# An evaluation of range accuracy in the Super Dual Auroral Radar Network over-the-horizon HF radar systems

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Abstract. The Super Dual Auroral Radar Network (SuperDARN) of over-the-horizon HF radars forms a powerful diagnostic of large-scale ionospheric and magnetospheric dynamics in the Northern and Southern Hemispheres. Currently, the ground location of the HF radar returns is routinely determined by a simple range-finding algorithm, which takes no account of the prevailing HF propagation conditions. This is in spite of the fact that both direct E and F region backscatter and  $1^{1}_{2}$ -hop E and F region backscatter are commonly used in geophysical interpretation of the data. Here HF radar backscatter which has been artificially induced by the high-power RF facility (ionospheric heater) operated by the European Incoherent Scatter Scientific Association at Tromsø is used to provide a range calibration for the SuperDARN radars. The known ground range, the measured radar slant range, and the group path calculated by a ray-tracing simulation are compared. The standard algorithm for backscatter ground range location is found to be accurate to within 16 km and 60 km for direct and  $1^{1}_{2}$ -hop backscatter, respectively.

# 1. Introduction

The SuperDARN network [Greenwald et al., 1995] currently consists of eight over-the-horizon HF radars in the northern polar regions and six radars in the southern polar regions. Radars of this design have been in operation since the deployment of the prototype Goose Bay radar in the early 1980s and currently form a powerful diagnostic of large-scale ionospheric and magnetospheric dynamics in the Northern and Southern Hemispheres. The radars are frequency-agile, operating between 8 and 20 MHz, and have fields of view which extend in range from 180 km to over 3000 km in standard operations. The radar systems rely on the refraction of the HF radiation both in order to achieve orthogonality to the Earth's magnetic field, a requirement for scattering off the ionospheric irregularities, which form the targets for such radar systems, and to achieve backscatter from the longer ranges, which require overthe-horizon operations. The ground locations of the HF radar returns are routinely determined by a simple rangefinding algorithm, which assumes straight line

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Paper number 2000RS002558. 0048-6604/01/2000RS002558\$11.00 propagation at the speed of light from the radar site to a target at a given altitude above the Earth and thus takes no direct account of the prevailing HF propagation conditions [*Baker et al.*, 1986]. In reality these propagation conditions are highly variable, with direct  $({}^{1}_{12}$ -hop) propagation to E and F region ionospheric irregularities and  $1{}^{1}_{12}$ -hop propagation to both the E and F region commonly observed by the radar systems. Previous studies have adopted a ray-tracing simulation [*Villain et al.*, 1984; *Baker et al.*, 1986] or velocity field cross-correlation [*Ruohoniemi et al.*, 1987; *André et al.*, 1997] approach to assess the accuracy of range-finding using the straight line approximation. These studies suggested agreement between the ground range and radar range within ~15 km over a  ${}^{1}_{2}$ -hop path.

Early studies with HF radars either used HF radar data alone or combined HF radar data with data from instruments such as ground magnetometers, which had a limited spatial resolution and are often only available from arrays which are sparsely populated in comparison to radar fields of view. However, the growing importance of combined ground-spacecraft measurements and multi-instrument studies from the ground have led to numerous coordinated studies with instruments of a high spatial resolution, such as meridian scanning photometers, all sky cameras, and auroral imagers. In these studies the location of the radar

Table 1. Geographic Locations of the Instrumentation

Instrument	Ground Location				
	Latitude, °N	Longitude, °E			
Þykkvibær	63.77	20.54			
Hankasalmi	62.32	-26.61			
Tromsø Heater (field aligned, 210 km)	69.20	19.2			

backscatter is a crucial element in the study. In addition, spacecraft overpasses, where the magnetic conjugate points are computed to a high precision using geomagnetic field models, require high accuracy in the location of the HF radar backscatter. A illustrative example is found in the recent studies made of the footprint of the magnetospheric cusp. Optical signatures associated with dayside reconnection have been extensively investigated from the ground in the visible wavelengths of 630.0 nm and 557.7 nm associated with auroral activity [e.g. Lockwood et al., 1993, 1995; Sandholt et al., 1996]. HF radars observe pulsed ionospheric flows in the cusp region, which are the response to this transient reconnection at the magnetopause [e.g. Pinnock et al., 1993, 1995; Provan et al., 1998, 1999]. Such events have been studied in detail by a variety of instruments [see e.g. Milan et al., 2000; Neudegg et al., 1999, and references therein]. The interpretation of such studies requires an accurate evaluation of radar range, in a situation where the radar backscatter may be either 1/2 hop or 11/2 hop [Milan et al. 1998].

In this paper, HF radar backscatter which has been artificially induced at a precisely known location by the high-power RF facility (ionospheric heater) operated by the European Incoherent Scatter (EISCAT) Scientific Association at Tromsø is used to provide a range calibration for the Super Dual Auroral Radar Network (SuperDARN) radars. The known ground range, the measured radar slant range, and the group path calculated from a ray-tracing simulation are compared. This demonstrates an excellent agreement between the measured radar slant range and the calculated group path and allows the quantification of the deduced range accuracy to be made for  $1/2^-$ ,  $1^{1}/2^-$  and  $2^{1}/2^-$ hop backscatter modes over paths of ~850 and ~1800 km.

# 2. Instrumentation

The data presented here result from the generation of

artificial ionospheric HF coherent backscatter. It is well known that the EISCAT heating facility at Tromsø, Norway [Rietveld et al., 1993], is capable of generating artificial field-aligned irregularities using high-power HF radio waves [e.g., Robinson, 1989]. These irregularities are detectable by both incoherent and coherent scatter radars [e.g., Robinson et al., 1997]. The Co-operative U.K. Twin Located Auroral Sounding System (CUTLASS) radar is a pair of HF coherent backscatter radar systems located at Hankasalmi, Finland, and Þykkvibær, Iceland, and forms part of the SuperDARN array. Details of SuperDARN are given by Greenwald et al. [1995] and details of the CUTLASS systems at Hankasalmi and Þykkvibær are given by Milan et al. [1997]. The CUTLASS radars' locations and the Tromsø heater's ionospheric interaction location are given in Table 1 and shown in geographic coordinates in Figure CUTLASS is ideally situated for making 1a. observations of the heated volume over Tromsø [e.g., Bond et al., 1997; Eglitis et al., 1998], as shown schematically in Figure 1b.

Here data are presented from the Observations of ULF waves with CUTLASS and the Heater (SP-UK-OUCH; Wright and Yeoman [1999]) experiment. In this experiment the CUTLASS radars run in a high-temporalresolution and high-spatial-resolution mode, with Hankasalmi running a six-beam scan (scanning beams 7 through 2, inclusive) with an integration time of 1 s, while Pykkvibær runs a three-beam scan (beams 13-15) with a 2 s integration period. Thus both radars produce data with a temporal resolution of 6 s. Both run in a high-spatial-resolution mode, with each radar having 75 range gates of 15 km length, centered on the heated volume at Tromsø (the distance to the first range gate being set at 480 and 1470 km for the Hankasalmi and Pykkvibær radars, respectively). The EISCAT heater was in continuous operation at 50% power (using  $6 \times 80$ kW transmitters, an Effective Radiated Power (ERP) of ~130 MW), at a frequency of ≈4-5 MHz for 4 hour intervals. Observations with the heater aligned along the magnetic field and employing a  $\pm 30^{\circ}$  beam swing at a frequency of 1 Hz, centered vertically, have been recorded, with the resulting backscatter detected by both radars of the CUTLASS system. The heater produces artificial electron density irregularities in the F region ionosphere, which act as targets for the HF radar. The artificial targets result in very high returned backscatter power in comparison to naturally occurring irregularities. This allows a short integration time to be run on the radar, providing higher time resolution than is normally available.



Figure 1. (a) The geographic locations of the Tromsø heater, and the fields of view of the CUTLASS Hankasalmi and Pykkvibær radars. (b) A schematic of the artificial backscatter experiment, SP-UK-OUCH. The Tromsø heater continuously heats the F region ionosphere, creating artificial ionospheric irregularities. These are detected by the CUTLASS radars, operating in a high temporal and spatial resolution mode.



**Plate 1.** Artificial backscatter generated in the Hankasalmi and Pykkvibær radars during the SP-UK-OUCH run of 1210-1250 UT, October 15, 1998, as a function of time and radar range gate. (a) Hankasalmi power, (b) Hankasalmi elevation angle, (c) Pykkvibær power, and (d) Pykkvibær elevation angle.

## 805

# 3. Observations

# 3.1. SP-UK-OUCH Data

Approximately 40 hours of the SP-UK-OUCH experiment have been run on the EISCAT ionospheric heating facility during 1998 and 1999. Ionospheric conditions are not always suitable for heating, however, and the radar modes run are also not always suitable for statistical analysis. Here we present data from 21 hours and 30 min of Hankasalmi radar backscatter and 12 hours and 40 min of Pykkvibær radar backscatter for which strong radar returns were obtained over a number of radar range gates for periods of greater than 1 hour. In addition to strong returns we also require that the radar was operating consistently at one sounding frequency. Data from an example run of SP-UK-OUCH will be presented in detail, along with statistical analysis of the entire data set, from both radar systems.

Plate 1 presents data from a section of an SP-UK-OUCH run from 1210 to 1250 UT on October 15, 1998, an interval when the heater was operating in its beam swinging mode. Returned backscatter power and elevation angle are plotted as a function of time and range gate for the Hankasalmi radar beam 5 (Plates 1a and 1b) and Pykkyibær radar beam 15 (Plates 1c and 1d) which overlie the heater location. During this interval the Hankasalmi radar was operating between 19.415 and 19.680 MHz, and the Pykkvibær radar was operating bêtween 12.105 and 12.235 MHz, both frequency bands having been selected at the beginning of the experiment to optimize the returned backscatter power. In Plate 1a a broad band of high-power radar returns, generated by the Tromsø heater, can be seen between range gates 25 and 35. Lower power backscatter is also visible beyond range gate 50, but these data have been identified by the radar, via their Doppler velocity and spectral width, as ground scatter: radiowaves that have refracted in the ionosphere and backscattered from the ground. The elevation angles for the heater-induced scatter (Plate 1b) are centered at 10°, with the nearer ranges exhibiting slightly higher elevation angles than the farther ranges, as expected for a constant irregularity height. The backscattered power from Pykkvibær is displayed in Plate 1c. In this case a threshold of 6 dB has been applied to the data, in order to reduce clutter from lowpower ground backscatter returns. Again, a band of high backscatter power is visible, this time between range gates 33 and 38. The Pykkvibær data for this interval are not as clear as the Hankasalmi data, and the unstructured low-power returns surrounding the artificial backscatter region are due to ground scatter. A second structured region of backscatter is visible centered on range gate 56. The elevation angles for the two structured regions of backscatter (Plate 1d) are centered on 18° and 31° for the nearer range and farther range regions, respectively. The origins of these two regions of backscatter will be discussed below in Sections 4.2 and 4.3.

## 3.2. Hankasalmi Radar Performance

In order to assess the performance of the radar system range evaluation the data from the interval presented in Section 3.1 and the overall SP-UK-OUCH data set from beam 5 of Hankasalmi and beam 15 of Pykkvibær require a more statistical analysis. In order to assess the overall performance and consistency of the radars the occurrence of backscatter with a power of >10 dB has been quantified as a function of the geographic latitude and longitude determined by the radar range-finding algorithm. Figure 2 presents the results of such an analysis for the Hankasalmi radar. The occurrence of radar backscatter at power levels above 10 dB is plotted as a function of geographic latitude derived from the standard radar rangefinding algorithm, with occurrence denoted on the left hand y axis labels. Also plotted is the actual latitude of the Tromsø field line as a function of altitude, denoted on the right hand y axis labels, taking into account the dip angle of the geomagnetic field. Figure 2a presents the Hankasalmi radar occurrence statistics for the interval displayed in Plate 1. The region of high-power artificial scatter occupies a broad latitude range (as determined by the standard range finding algorithm) between 68.5°N and 70°. In this case the range-finding algorithm is used with a fixed height of 210 km, which is an estimate of the heater interaction height (the upper hybrid height, which lies between the X- and O-mode reflection heights, Robinson [1989]), as determined from the heater frequency and measurements from the EISCAT dynasonde at Tromsø. It is informative to compare such occurrence statistics with similar experiments run with a field-aligned heater. Such an experiment was run on October 18 1999, with the Hankasalmi radar again operating at 19 MHz, and is displayed in Figure 2b. The field-aligned results show a narrower band of artificial backscatter, as expected, with the center of the region lying at 69.1°N. Results from the Hankasalmi radar are, in fact, highly consistent, and the entire data set is plotted in Figure 2c, again showing a peak occurrence at 69.1°N, where almost 7000 data points have been accumulated. The width of the distribution in latitude in this case is some 30% larger than that in the case study patch width presented in Figure 2b.







Figure 3. Occurrence rates for artificial backscatter detection at a power of >10 dB for Þykkvibær, as a function of radar-deduced longitude. The location of the Tromsø field line ground longitude (the true location of the center of the artificial irregularity patch) is also shown as a solid line. (a) the SP-UK-OUCH run of 1210-1250 UT, October 15, 1998, as shown in Plate 1; (b) the entire 12 hours and 40 min of SP-UK-OUCH data analyzed from Þykkvibær for the years 1998 and 1999.

## 3.3. Þykkvibær Radar Performance

Statistical results from the Pykkvibær radar are displayed in Figure 3. Here we do not distinguish between beam swung and field-aligned operations of the heater, as the heater beam swinging is limited to being near-orthogonal to the look direction of beam 15 of Pykkvibær. The occurrence of radar backscatter at power levels above 10 dB is plotted as a function of geographic longitude derived from the standard radar range finding algorithm, and the actual longitude of the Tromsø field line is also indicated. The field is near orthogonal to the radar look direction, so the altitude variation of the field line position does not need to be considered in this case. Figure 3a presents the Pykkvibær radar occurrence statistics for the interval displayed in Plate 1. The ground scatter clutter is visible, along with two distinct regions of structured artificial



Figure 4. Range h as a function of sounding frequency f measured by the EISCAT dynasonde at Tromsø at 1248 UT on October 15, 1998. O-mode echoes are plotted as grey circles, and X-mode echoes are plotted as solid circles. The solid black curve indicates the real height profile as calculated by the Polan inversion algorithm. The profile parameters are indicated in the top left-hand corner of the plot. The dashed curve illustrates the modeled double-Chapman-layer profile utilized for the ray tracing presented in this paper. The model ionospheric parameters are given in Table 2.

backscatter, the main heater signature being located at apparent longitudes of  $22^{\circ}-23^{\circ}E$ , and the farther range scatter being located at apparent longitudes of  $29^{\circ}-30^{\circ}E$ . The entire data set is plotted in Figure 3b, showing a peak occurrence at  $23^{\circ}E$ , where over 3600 datapoints have been accumulated.

#### 3.4. Ray-tracing Simulation

In order to gain a more detailed understanding of the propagation conditions prevailing during the experiment of October 15 1998, a ray-tracing simulation has been undertaken. A modified version of the ray-tracing code developed by *Jones and Stephenson* [1975] has been used in which the angle between the k vector and the magnetic field has been used to identify regions where backscatter is likely to occur in the presence of density irregularities. An example ionogram from the Tromsø dynasonde is illustrated in Figure 4, together with a real height profile, as calculated by the Polan inversion algorithm, and the two- $\alpha$ -Chapman-layer model which was used in the ray-tracing simulation. The critical frequencies of the two layers have been determined from the ionosonde measurements, and the ionospheric parameters for the two  $\alpha$ -Chapman layers employed in the simulation are given in Table 2. The critical and scale heights of the layers have also been chosen to be in

Region	Critical Frequency, MHz	Critical Height, km	Scale Height, km	
E	2.0	110	10	
F	8.5	244	35	

Table 2. Ionospheric Parameters Used in the Ray Tracing Simulation

broad agreement with the values returned by the International Reference Ionosphere (IRI) [Bilitza, 1990]. It should be noted that although we have assumed that the density profile remains fixed with position, this assumption is not unreasonable since the particular interval under study lies around local noon, and the values of  $f_0F_2$  observed by the Tromsø ionosonde are remarkably stable for the entire period. Launch angles for the rays were chosen within 1° of the mean elevation angles measured by the radars, with ray-trace frequencies of 19.5 MHz chosen for Hankasalmi and 12.2 MHz for Þykkvibær, near the centers of the narrow radaroperating-frequency bands. It was assumed that the backscattered power would return to the radar on the same path as on the outward leg (i.e., reciprocity is assumed to hold). The results obtained from the raytracing simulation of three propagation paths (one from Hankasalmi-Tromsø and two from Þykkvibær-Tromsø) are presented in Plate 2. For the case of the Hankasalmi radar, orthogonality occurs just prior to reflection while for the Þykkvibær radar orthogonality takes place close to or at the reflection point.

# 4. Discussion

Artificial coherent HF radar backscatter generated by the EISCAT ionospheric heating facility at Tromsø has provided a high-power signature at a known ground range which can be used to calibrate the range assessment of the SuperDARN radar facilities. Case study and statistical analysis of runs of the SP-UK-OUCH experiment have provided a database to assess both the accuracy and variability of the standard SuperDARN range-finding algorithm. Here the data presented above will be compared with ray-tracing results to assess the propagation paths concerned. Three regions of structured artificial backscatter (one measured by the Hankasalmi radar and two measured by the Þykkvibær radar) are analyzed in terms of three propagation paths: direct ionospheric backscatter from the Hankasalmi-Tromsø path (hereafter referred to as the  $1_{2}$ -hop path), ionospheric backscatter from a  $1_{2}$ -hop mode across the Pykkvibær-Tromsø path (hereafter referred to as the  $1^{1}_{2}$ -hop path), and ionospheric backscatter from a  $2^{1}_{2}$ -hop mode across the Pykkvibær-Tromsø path (hereafter referred to as the  $2^{1}_{2}$ -hop path).

#### 4.1. Half-Hop Path Propagation

The artificial scatter observed from the Hankasalmi radar presented in Plate 1 and Figure 2 resulted from a radar frequency of 19 MHz, with a heater interaction height of 210 km and with radar elevation angles of  $\sim 10^{\circ}$ . A ray trace for this situation is presented in Plate 2a. The HF rays achieve the orthogonality required for HF backscatter at ground ranges of ~850 km, prior to reflection, and hence the propagation path for this scatter can be determined with confidence. The ground range, the calculated group path, and the measured radar slant range for the three propagation paths are summarized in Table 3. The field-aligned data from Figure 2b are used for this comparison, as this is the most appropriate mode for this purpose, with the heater beam aligned with the Tromsø field line. In this case the strongest heating effects (due to field-aligned heating [Robinson, 1989]) will occur in a very well defined location. The ray trace reveals a difference between the ground range and the group path for this mode of 50 km and a difference between the measured radar slant range and the calculated group path of only 1.6 km (radar slant range here has a resolution of 15 km). The range-finding algorithm then calculates the backscatter ground latitude at 69.06°. This agrees with the known ground latitude of the Tromsø heater interaction height at 210 km altitude within 16 km, in practice an accuracy of one 15 km range gate, well within the 45 km range gates used in standard radar operations. Comparing the field-aligned heater measurements in Figure 2b with the measurements resulting from heater beam-swinging operations in Figure 2a, the beam swinging can be seen to result in a broader patch of heater-induced irregularities, which are centered between a location vertically above the heater (the center of the heater beam swinging) and the field-aligned position where heating is at its most effective.





**Plate 2.** Ray traces [*Jones and Stephenson*, 1975] for the three HF propagation paths under analysis: (a) the  $1_{T}$  hop path from Hankasalmi to Tromsø at elevation angles 9°-11°; (b) the  $1_{T}$  hop path from Pykkvibær to Tromsø at elevation angles 17°-19°; and (c) the  $2_{T}$  hop path from Pykkvibær to Tromsø at elevation angles 30°-32°. The location of the Tromsø file line is also indicated in each panel. Ray tracing was calculated for frequencies of 19.5 MHz for Hankasalmi and 12.2 MHz for Pykkvibær, near the centers of the narrow radar-operating-frequency bands.

distance (km)

500

6400

Path	Raytrace Height, km	Ground Range (GR), km	Calculated Group Path (GP), km	Measured Slant Range (SR), km	GP-GR, km	SR-GP, km	Radar-Deduced Lat./ Long.	Radar Range Error, km
Hankasalmi, 10° elevation	198	825.4	875 4	877	50	1.6	69.06° ± 0.13°	16 ± 15
Pykkvibær, 18° elevation	150	1829	1986	1987	157	1	22.2° ± 0.4°	114 ± 15
Þykkvibær, 31° elevation	199	1825	2205	2272	380	67	29.5° ± 0.4°	390 ± 15

Table 3. Details of the Calculated and Measured Ray Paths to the center of the Artificial Scatter Region at Tromsø, at an Assumed Radar Backscatter Altitude of 210 km

#### 4.2. One-and-One-Half-Hop Path Propagation

The main band of artificial scatter observed from the Þykkvibær radar presented in Plate 1 and Figure 3 resulted from a radar frequency of 12 MHz, with radar elevation angles of ~18°. A ray trace for this situation is presented in Plate 2b. The HF rays achieve orthogonality at ground ranges of ~1800 km, and thus the propagation path for this scatter can again be determined with confidence (the rays also achieve orthogonality at their 1/2-hop position, but this lies outside the radar field of view in this case). The parameters deduced for this propagation path and the data presented in Figure 3a are summarized in Table 3. The raytrace reveals a difference between the ground range and the calculated group path for this mode of 157 km and a difference between the measured radar slant range and the group path of only 1 km. The rangefinding algorithm then calculates the backscatter ground longitude at 22.2°. This differs from the known ground longitude of the Tromsø heater by 114 km, an accuracy of seven 15 km range gates, or just over two 45 km range gates as used in standard radar operations.

## 4.3. Two-and-One-Half-Hop Path Propagation

The second population of artificial scatter observed from the Pykkvibær radar at farther range gates was observed with radar elevation angles of  $\sim 31^{\circ}$ . A ray trace for this situation is presented in Plate 2c. The HF rays again achieve orthogonality at ground ranges of  $\sim 1800$  km, and thus the propagation path for this scatter can again be determined with confidence. The parameters deduced for this propagation path are summarized in Table 3. The ray trace reveals a difference between the ground range and the calculated group path for this mode of 380 km and a difference between the measured radar slant range and the calculated group path of 67 km. The latter discrepancy results from accumulated errors in the ray trace calculation and inaccuracies in the (uniform) ionospheric models employed. The range-finding algorithm then calculates the backscatter ground longitude at 29.5°. This differs from the known ground longitude of the Tromsø heater by 390 km, an accuracy of twenty-six 15 km range gates, or eight 45 km range gates, as used in standard radar operations.

#### 4.4. Implications for Radar Operations

The comparison presented in Sections 4.1-4.3 shows an excellent agreement between the calculated ray trace results and the measured radar slant range. This gives confidence in the accuracy of the model ionospheres employed and the propagation modes inferred. It also demonstrates the accuracy of the radar elevation angle determination. The comparison between the known target ground location and the ground location inferred from the range-finding algorithm can then be used to evaluate the performance of the algorithm under more general conditions. For the 1/2-hop path over a ground range of 825 km the range-finding algorithm has an accuracy of well within the standard 45 km range gate of a SuperDARN radar. In fact, the accuracy of the range finding over such a path is close to the narrowest range gate such systems are capable of at present. The HF propagation to the point of backscatter under these conditions at 19 MHz is close to straight line propagation, so this result is perhaps unsurprising. The agreement between the ground range and radar range within ~15 km is fully in accord with the results of Ruohoniemi et al. [1987], who employed a ray-tracing analysis and a velocity field cross-correlation technique between the Goose Bay HF radar and the Sondrestrom incoherent scatter facility for a 1500 km <sup>1</sup>/<sub>2</sub>-hop path. It also confirms the results of André et al. [1997], who performed a similar velocity field analysis to determine the relative accuracies of various HF radar frequencies

over a  ${}^{1}_{2}$ -hop path, again finding an accuracy of 15 km. André et al. [1997] suggested an optimum assumed height for  ${}^{1}_{2}$ -hop propagation paths of 200-300 km, which is in accord with the results presented here.

For the  $1^{1}/_{2}$ -hop path over 1830 km an overestimate of the radar range of ~100 km is deduced as a consequence of the group path of the HF rays exceeding that of the straight line propagation assumed in the range-finding algorithm. In the standard use of the range-finding algorithm, a height of 300-400 km is assumed to offset the expected difference between the group path and ground range. This altitude leads to a reduced range offset, 60 km. The results presented here suggest that an additional offset of ~60 km (or an assumed height of 500 km) would improve the performance further for  $1^{1}/_{2}$ -hop backscatter.

The  $2^{1}/_{2}$ -hop path exhibits even more significant range errors, with the range-finding algorithm overestimating the length of the 1830 km path by 390 km. In contrast to the  $1^{1/2}$ -hop path no simple adjustment of the assumed backscatter height can be employed to compensate for the range offset, with the adoption of an assumed target altitude of 500 km only reducing the range offset to 300 km. Two-and-one-half-hop paths are not observed in natural scatter, because of their low power, so this is not in practice a problem for standard radar operations. It does, however, suggest that care should be taken in the use of high-elevation-angle scatter. Baker and Greenwald [1988] demonstrated a relative range error of 120 km between 15° and >20° elevation angle backscatter, with ionospheric tilts thought to be responsible for the different propagation paths observed. The artificial backscatter data presented in Figures 2 and 3 are extremely consistent, suggesting that the range offsets determined here are highly systematic. No detectable variation is seen in the location of the artificial backscatter in the Hankasalmi field of view, over the 850 km,  $1_{2}$ -hop path. A variability of only two 15 km range gates is seen in the 1830 km,  $1^{1}_{2}$ -hop path from Þykkvibær to Tromsø. This stability in the measured range gates of the artificial scatter persists even when the location of naturally occurring ground backscatter varies by at least 20 range gates (300 km) during 4 hour runs of SP-UK-OUCH, indicating changes in the prevailing ionospheric conditions. This observation offers experimental confirmation of the modeling of André et al. [1997], who suggested a group path change of <10 km due to a change in the altitude of the F region peak by 50 km. It also demonstrates that the accuracy in radar range is maintained to within 30 km for  $1^{1/2}$ -hop paths. Thus the accuracy in the determination of the location of the backscatter is only weakly controlled by variations in the ionosphere between the radar and target ionosphere. Under most conditions the use of a full ray tracing analysis based on ionosonde observations or on predictive models of the ionosphere is not required for the range accuracy needed for geophysical interpretation of HF radar data.

# 5. Summary

An evaluation of the absolute range finding accuracy of current routine analysis of the SuperDARN network of over-the-horizon HF radars has been performed, comparing ground range, calculated group path, and measured radar slant range of backscatter artificially excited at a known location by the EISCAT heating facility at Tromsø. HF propagation over a 1/2-hop path, a  $1^{1}/_{2}$ -hop path, and a  $2^{1}/_{2}$ -hop path has been examined for the first time. The radar slant range and the calculated group paths are in excellent agreement for all three paths. The standard algorithm for backscatter ground range location is found to be accurate to within 16 km and 114 km for  $\frac{1}{2}$ -hop and  $\frac{1}{2}$ -hop backscatter, respectively, when using the true backscatter height, and these range offsets are extremely consistent. Using an assumed backscatter height of 200-300 km for <sup>1</sup>/<sub>2</sub>-hop paths thus gives an accurate range determination, to within ~15 km, as suggested by André et al. [1997]. The standard assumption of 400 km height for far range backscatter reduces the error for 1<sup>1</sup>/<sub>2</sub>-hop backscatter to 60 km, but the adoption of an additional range offset of 60 km (or the adoption of an assumed backscatter altitude of 500 km) seems desirable for 1<sup>1</sup>/<sub>2</sub>-hop paths. Two-and-onehalf-hop paths, although not seen in practice, would introduce significant range errors, and suggest that highelevation-angle backscatter should be interpreted with caution.

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# References

André, R., C. Hanuise, J.-P. Villain, and J.-C. Cerisier, HF radars: Multifrequency study of refraction effects and localization of scattering, Radio Sci., 32, 153-168, 1997.

- Baker, K. B., and R. A. Greenwald, The vertical angle of arrival of high-frequency signals propagating from Thule to Goose bay, *Johns Hopkins APL Tech. Dig.*, 9, 121-130, 1988.
- Baker, K. B., R. A. Greenwald, A. D. M. Walker, P. F. Bythrow, L. J. Zanetti, T. A. Potemra, D. A. Hardy, F. J. Rich, and C. L. Rino, A case study of plasma processes in the dayside cleft, *J. Geophys. Res.*, 91, 3130-3144, 1986.
- Bilitza, D. (Ed.), International Reference Ionosphere 1990, NSSDC 90-22, http://nssdc.gsfc.nasa.gov, Natl. Space Sci. Data Cent., Greenbelt, Md., 1990.
- Bond, G. E., T. R. Robinson, P. Eglitis, D. M. Wright, A. J. Stocker, M. T. Rietveld and T. B. Jones, Spatial observations by the CUTLASS coherent scatter radar of ionospheric modification by high power radio waves, Ann. Geophysicae, 15, 1412-1421, 1997.
- Eglitis, P., T. R. Robinson, M. T. Rietveld, D. M. Wright and G. E. Bond, The phase speed of artificial irregularities observed by CUTLASS during HF modification of the auroral ionosphere, J. Geophys. Res., 103, 2253-2259, 1998.
- Greenwald, R. A., et al., DARN/SuperDARN A global view of the dynamics of high-latitude convection, *Space Sci. Rev.*, 71, 761-796, 1995.
- Jones, R. M., and J. J. Stephenson, A versatile threedimensional ray tracing computer program for radio waves in the ionosphere, OT 75-76, Off. of Telecommun., U.S. Dep. of Comm., Washington, D.C., 1975.
- Lockwood, M., J. Moen, S.W.H. Cowley, A.D. Farmer, U.P. Løvhaug, H. Lühr, and V.N. Davda, Variability of dayside convection and motions of the cusp/cleft aurora, *Geophys. Res. Lett.*, 20, 1011-1014, 1993.
- Lockwood, M., S.W.H. Cowley; P.E. Sandholt, and U.P Løvhaug, Causes of plasma flow bursts and dayside auroral transients: An evaluation of two models invoking reconnection pulses and changes in the Y component of the magnetosheath field, J. Geophys. Res., 100, 7613-7626, 1995.
- Milan, S.E., T.K. Yeoman, M. Lester, E.C. Thomas, and T.B. Jones, Initial backscatter occurrence statistics from the CUTLASS HF radars, Ann. Geophysicae, 15, 703-718, 1997.
- Milan, S. E., T. K. Yeoman, and M. Lester, The dayside auroral zone as a hard target for coherent HF radars, *Geophys. Res. Lett.*, 25, 3717-3720, 1998.
- Milan, S.E., M. Lester, S.W.H. Cowley and M. Brittnacher, Convection and auroral response to a southward turning of the IMF: Polar UVI, CUTLASS, and IMAGE signatures of transient magnetic flux transfer at the magnetopause, J. Geophys. Res., 105, 15,741-15,755, 2000.
- Neudegg, D.A., T.K. Yeoman, S.W.H. Cowley, G. Provan, G. Haerendel, W. Baumjohann, U. Auster, K.-H. Fornacon, E. Georgescu, and C. J. Owen, A Flux transfer event observed at the magnetopause by the Equator-S spacecraft and in the

ionosphere by the CUTLASS HF radar, Ann. Geophysicae, 17, 707-711, 1999.

- Pinnock, M., A.S. Rodger, J.R. Dudeney, K.B. Baker, P.T. Newell, R.A. Greenwald, and M. Greenspan, Observations of an enhanced convection channel in the cusp ionosphere, J. Geophys. Res., 98, 3767-3776, 1993.
- Pinnock, M., A.S. Rodger, J.R. Dudeney, F. Rich, and K.B. Baker, High spatial and temporal resolution observations of the ionospheric cusp, Ann. Geophysicae, 13, 919-925, 1995.
- Provan, G., T.K. Yeoman, and S.E. Mılan, CUTLASS Finland radar observations of the ionospheric signatures of flux transfer events and the resulting plasma flows, Ann. Geophysicae, 16, 1411-1422, 1998.
- Provan, G., T.K. Yeoman, and S.W.H. Cowley, The influence of the IMF B<sub>y</sub> component on the location of pulsed flows in the dayside ionosphere observed by an HF radar, *Geophys. Res. Lett.*, 26, 521-524, 1999.
- Rietveld, M. T., H. Kohl, H. Kopka and P. Stubbe, Introduction to ionospheric heating at Tromsø, I, Experimental overview, J. Atmos. Terr. Phys., 55, 577-599, 1993.
- Robinson, T. R., The heating of the high latitude ionosphere by high power radio waves, *Phys. Rep.*, 179, 79-209, 1989.
- Robinson, T. R., A. J. Stocker, G. E. Bond, P. Eglitis, D. M. Wright and T. B. Jones, O- and X-mode heating effects observed simultaneously with the CUTLASS and EISCAT radars and low power HF diagnostics at Tromsø, Ann. Geophysicae, 15, 134-136, 1997.
- Ruohoniemi, J. M., R. A. Greenwald, K. B. Baker, J.-P. Villain, and M. A. McCready, Drift motions of small-scale irregularities in the high-latitude F region: An experimental comparison with plasma drift motions, J. Geophys. Res., 92, 4553-4564, 1987.
- Sandholt, P.E., C.J. Farrugia, P. Stauning, S.W.H. Cowley, and T. Hansen, Cusp/cleft auroral forms and activities in relation to ionospheric convection: Responses to specific changes in solar wind and interplanetary magnetic field conditions, J. Geophys. Res., 101, 5003-5020, 1996.
- Villain, J.-P., R. A. Greenwald, and J. F. Vickrey, HF ray tracing at high latitudes using measured meridional electron density distributions, *Radio Sci.*, 19, 359-374, 1984.
- Wright, D. M., and T. K. Yeoman, High resolution bistatic HF radar observations of ULF waves in artificially generated backscatter, *Geophys. Res. Lett.*, 26, 2825-2828, 1999.

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