

1 **DR2015305**

2 **SUPPLEMENTARY MATERIAL**

3 Appendix A: Detailed description of methods used for geochemical modeling and list of
4 supplementary references.

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6 DR1: Whole-rock analyses of biotite and calcic A-type granites, with estimated magmatic
7 parameters and degree of partial melting.

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9 DR2: Partition coefficients for selected trace elements used for geochemical modeling.

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11 DR3: Petrographic description of Neoproterozoic intermediate rocks of the Cobequid Highlands.

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13 DR4: Whole-rock analyses of intermediate and mafic rocks from the Cobequid Highlands used
14 as sources for geochemical models.

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APPENDIX A

Methodology for geochemical models

Fractionation of mafic sources

The fractionation models used in this work include a two-step process. First the changes in major elements through fractionation were modeled by the use of PELE software for Windows (Boudreau, 1999). Whole-rock analyses of selected mafic sources were used as input. PELE uses a list of mineral phases that can crystallize during the fractionation sequence. From the given list certain minerals were excluded such as leucite, nepheline, mellilite, garnet, kalsilite, whitlockite, corundum, perovskite, muscovite, troilite, pyrrhotite, pyrite, sulphide liquid, graphite, calcite, dolomite, anhydrite and immiscible liquid. Although pyrite is reported as an accessory phase in the Wentworth gabbros, the exclusion of this mineral is not considered to significantly compromise the model, since there no S was determined in whole-rock for these rocks, and therefore there was no input for that element. Other parameters involve temperature, pressure, oxygen fugacity, entropy, enthalpy and volume. In the absence of constraints for entropy, enthalpy and volume, an isobaric fractionation was assumed and was tested for a range of pressures from 2 to 0.6 GPa. An initial 0.5 wt.% H₂O was assumed although not measured. An anhydrous composition for the Fountain Lake basalts is assumed from their mineralogy which includes olivine, plagioclase, pyroxene and Fe-oxides. Fractionation models were performed under different oxygen fugacities ranging from NNO to +3 QFM. For different oxygen fugacities the program will adjust the amounts of FeO and Fe₂O₃ in the whole-rock input. The range of temperature for all runs was from the liquidus, as determined by the program for a given pressure

and oxygen fugacity, to 800 °C, which is an average zircon-saturation temperature for the studied granites.

After the parameters were set, PELE performed a fractionation simulation until the lowest temperature was achieved. During the simulation, the program predicted mineralogical assemblages that would be in equilibrium with the remaining melt for a given temperature. For each equilibrium stage (or “step” as reported by the program) the following are predicted: mass of liquid, mass of solids, fractionating mineral phases (chemistry and mass for each phase), whole-rock chemistry of the remaining melt, viscosity, volume, enthalpy, entropy and oxygen fugacity. Any invalid parameters or unrealistic set of parameters would crash the program.

The most realistic output for the fractionation models was considered to be from the run at 0.6 GPa, under an oxygen fugacity of -2 QFM. The fractionated minerals predicted were olivine, plagioclase, pyroxene, apatite, and Fe-Ti oxides. The minerals predicted to remain in the melt at the latest stages were alkali feldspar, quartz, titanite, ilmenite, and aenigmatite. Since all the predicted fractionating minerals reflect the mineralogy of the mafic rocks, the output was considered reliable. The predicted viscosity was comparable with the viscosity estimates for the sodic granites, using the method of Giordano et al. (2008). The felsic melt under these conditions was produced in ten steps.

As a second procedure the trace elements were modeled using Rayleigh fractionation (Rayleigh, 1896) as applied by Neuman et al. (1954) which is described by the equation

$$C_L = C_O F^{(D-1)} \quad (1)$$

Where C_L : concentration of element in remaining liquid, C_O : concentration of element in the parent melt, F : fraction of remaining melt defined as M_L/M_O (M_L : mass of original melt, M_O :

mass of remaining melt), D: the bulk partition coefficient for element calculated as $\sum X_A D_i$ (X_A : weight fraction of mineral A in the rock, D_i : partition coefficient of element i in mineral A).

Trace elements were therefore modeled for each step produced by PELE, where D was estimated by the modal amounts of the predicted fractionating minerals and the chemistry of the remaining melt, and F was calculated by the estimated masses of the original and remaining melts. Since the produced melts varied from mafic to felsic, the partition coefficients used, varied accordingly. For produced melts with SiO₂ up to 52 wt.% the partition coefficients used were for basaltic melts. For melts from 52 to 54 wt.% the partition coefficients used were those for basaltic andesite. For melts with 54 to 62 wt.% SiO₂ partition coefficients for andesite were used. For melts with SiO₂ ranging from 62 to 69 wt.% partition coefficients for dacites were used. Melts more silicic than 69 wt.% were considered rhyolitic and the appropriate partition coefficients were used (Table DR2).

Partial melting models

Batch partial melting was modeled by the equation

$$C_L/C_0 = 1/(D(1-F)+F) \quad (2)$$

Where C_L : the concentration of element in extracted melt, C_0 : concentration of element at the source, D: bulk partition coefficient as defined above, F: fraction of extracted melt. For highly incompatible elements such as Rb, Nb, and Th, the above equation may be simplified as:

$$C_L/C_0 = 1/F \quad (3)$$

which if rearranged will be:

$$F = C_0/C_L \quad (4)$$

Using the above simplified equation for selected highly incompatible elements, and their average concentrations in Neoproterozoic intermediate rocks, the degrees of partial melting were estimated. The rest of the trace elements were modeled with the original equation for the range of F values determined by the highly incompatible elements.

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